

Some Characteristics of Seismic Activity in the Song Tranh 2 Reservoir, Quang Nam, Vietnam by Local Seismic Network Data

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Abstract: The Song Tranh 2 hydropower construction is located in the Quang Nam province (central Vietnam), it has a reservoir volume of 740 million cubic meters of water and a dam height of 96 m. The reservoir was filled to capacity for the first time in February 2011 to about 160m, then it dropped to 140 m in July 2011. The filling of reservoir started again in August 2011 and the maximum water level of 175 m was reached in October 2011. Song Tranh 2 and its surrounding regions suffered from earthquakes in March 2011, it also suffered from a higher magnitude earthquake in October and November, 2012 of M=4.6 and 4.7 respectively. By the end of 2012, a seismic network including 6 stations was set up around the Song Tranh 2 reservoir area and a full network of 10 stations was set up by August 2013. The final seismic network is capable of detecting and locating weaker seismic earthquakes. It is also possible to calculate the extended source parameters like focal mechanism, slip direction etc. In the period from August 2013 to May 2014 about 2000 seismic events were detected, and 359 of them were localized and magnitudes were calculated. Seismic analysis and update catalogues are currently being conducted. The LocSAT application was used to locate events, to perform hypocentral inversion of the phase arrival data, to estimate the origin time, epicentral location, and depth by registered data from the VERIS network. Mechanism solutions with P-wave amplitude inversion of three events were determined. It suggests that reservoir construction is a major factor in the seismogenic process.

Keywords: Reservoir Induced Seismicity, Seismic Network, Source Parameters, Focal Mechanism, Completed Catalogue

1. Introduction

Earthquakes can be caused by natural processes (tectonic movement, volcanic activity etc.) or by human activities. The earthquakes associated with human activities are called seismic stimulation. Under certain suitable geological conditions anthropogenic activity can trigger or induce earthquakes [1]. The triggered or induced earthquakes are known to have occurred due to gold and coal mining [18], petroleum production, filling of artificial water reservoirs, high-pressure liquid injection into ground and natural gas production. Induced seismicity includes cases such as seismicity associated with water reservoirs, underground mining, large-scale surface quarrying, high-pressure fluid injection for geothermal power generation, oil production,

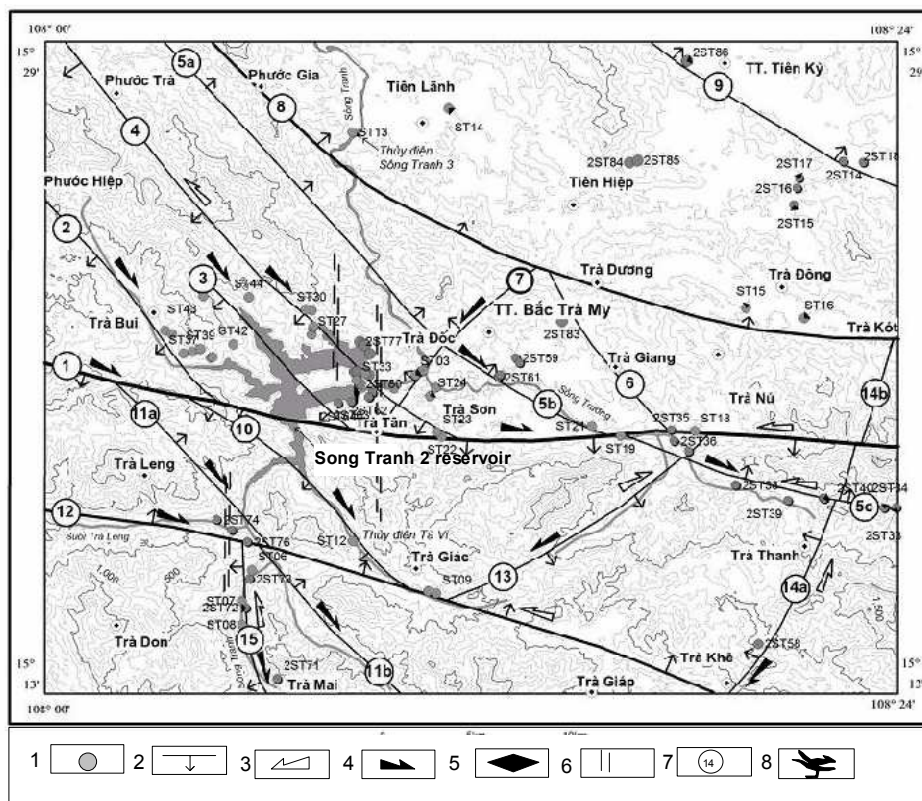
solution mining and waste disposal, removal of underground fluids such as oil, gas, water, and underground explosions [22].

In some locations the Reservoir Induced Seismicity (RIS) occurs early after reservoir impoundment, while in other places the earthquake excitation occurs after a number of years when the reservoir water began to rebuild. Based on such observations, Simpson (1986) and Gupta (1992) have classified the RIS into two types: fast response type and slow response type [12, 22]. Fast response-type is directly calculated increase in seismic analysis phase of the first water reservoir, or a sudden change in the seismic properties due to rapid changes in reservoir water levels. Examples of such reactions are the Lake Monticello, Lake Manic 3, Nurek and Kariba reservoirs. Nurek has had apparent rapid response to change after the reservoir water level. In contrast with the

The largest reservoir impoundment - triggered earthquakes exceeded magnitude 6, while the largest earthquakes triggered by the other engineering activities was around magnitude 5. There are more than 70 examples of reservoir-induced seismicity (RIS) in the world's [11] associated earthquakes that impounded artificial water reservoirs, this was pointed out for the first time by Carder (1945) for Lake Mead in the United States [6]. India recorded a 6.3 magnitude earthquake at Koyna Dam, in December 1967. The number of reported cases of reservoir-induced earthquakes has steadily grown since then [12]. Hoa Binh reservoir, Northwest part of Vietnam, is a case of reservoir triggered seismicity where in 1989 was the largest earthquake of a magnitude of 4.9 [11]. This paper considers a case of reservoir triggered seismicity in the Song Tranh 2 reservoir located in the Quang Nam province in Central Vietnam. The Song Tranh 2 reservoir was filled to capacity for the first time in November 2010. The largest earthquake of a

On the basis of co-operation research between IGPVAST and IGFPAS which was signed in 2013, the installation of seismic stations, observation, seismic monitoring and processing, analysis and interpretation of the seismological data of reservoir Song Tranh 2 area in Bac Tra My district and surrounding areas in Quang Nam province in Central Vietnam were carried out. Special emphasis was put on investigating the mechanism of eruptions and seismic hazard estimation. The first results of the monitoring activity are presented below with the preliminary interpretation of seismic pattern.

The Song Tranh 2 (ST2) dam is one of a few dams planned to be set up on the River Tranh. Geographic co-ordinates of the dam are: 108° 08,832' E and 15°20,013' N. The reservoir volume is 740 million cubic meters of water and its height is 196 m (Fig.1).



Characteristics of geological background hydroelectric reservoir area ST2 are well developed and published [14, 16, 24]. The original rock in the area studied is primarily a metamorphic amphibolite general, formed during orogenic

movement Indosini. It happened about 245±250 million years and the granite slab was not [21]. The lamina of the metamorphic rock had extended dominant direction in sub-latitude, double parallel spot was winding and steep angles of

plug from 45° to descent. The ST2 was built on the metamorphic rocks of the Tra Bong and Kham Duc complex formations. Research has shown that the research area and its surroundings have undergone several stages of tectonic movements, with different distortional characteristics. There were two main ductile deformational phases with the general movement mechanism slipping by, including: early phase axis compressive strain axis, sub-latitude. Under this stress field, the NW-SE fracture activity slid to the left. In contrast, phase distortion of late had featured local compressive axial sub-meridian and enjoyed the strain axes of sub-latitude. Under the action of this stress field, the NW-SE fracture slid right; Cenozoic tectonic motion phases later had close ties with volcanic activities in Central and South Vietnam during the Neogene-Quaternary. The Dam area is located within a series of zones with split-snowboarding by NW-SE direction, which tended to raise the compression zone sub-meridian direction, which also creates a spreading trend E-W. The sub-latitude and NE-SW directions faults are underdeveloped while the sub-meridian direction fault is hardly absent (Fig. 1). The Dam is placed completely on the metamorphic complex

gneiss Kham Duc-Nui Vu of regional metamorphism originated in the Ordovician.

3. Seismic Monitoring of the Region of Dam ST2

We can divide the recording of seismicity in ST2 area into three periods (stages).

3.1. Period I - from January 2011 to October 2012

In the beginning of filling the ST2 reservoir, the area was monitored by two seismic stations in Binh Dinh and Hue, but staying away from the reservoir. They recorded the earthquake in the region with a magnitude of $M > 2.0$. From January 2011 to September 2012, the two stations recorded more than 100 quakes in the Bac Tra My and surroundings, having magnitudes of $M 1.8 \div 4.2$. The basic parameters of the two seismic stations in Binh Dinh and Hue are presented in Tab. 1.

Tab. 1. The location and equipment for seismic stations in period I.

No.	Station name	Code	Co-ordination		Elevation [m]	Logger / Seismometer
			$\Phi [^\circ\text{N}]$	$\lambda [^\circ\text{E}]$		
1	Binh Dinh	BDVB	13.8645	109.1111	61	Q330HRS / STS-2
2	Hue	HUBV	16.4155	107.5689	20	Q330 / Trillium-40

3.2. Period II - from October 2012 to August 2013

The monitoring of the induced seismicity in the ST2 reservoir, from October to November 2012, the seismographic network of five stations was deployed by IGP VAST in the ST2 reservoir and vicinity. The name of the stations were Tra Doc, Tra Bui, Tra Nu, Tra Mai and Tien Lanh were also used in period III (Tab. 2). They were equipped in period II with Guralp seismometers and the SAMTAC logger and one Trillium-40 seismometer with the Q330 logger (Tra Doc station). SAMTAC recorders did not have the possibility of data transmission online. After the installation of local seismic station networks, it was possible then to get a magnitude less than 2.0. After the seismic network station was set up there, it could record a magnitude of $M < 2.0$ upwards.

3.3. Period III - from August 2013 - New Joint Seismic Network

In August 2013, installed 10 seismic stations were installed.

Tab. 2. The location and equipment of local seismic stations of VERIS network (period III) for ST2 region (In brackets are loggers and seismometers used in period II).

No.	Station name	Code	Co-ordination		Elevation [m]	Logger / Seismometer
			$\Phi [^\circ\text{N}]$	$\lambda [^\circ\text{E}]$		
1	Tien Ngoc	TNG	15.4472	108.2038	97	NDL / LE-3Dlite
2	Tra Don	TDO	15.2432	108.0849	185	NDL / LE-3Dlite
3	Tra Leng	TLE	15.2722	108.0225	192	NDL / LE-3Dlite
4	Tra Giac	TGI	15.2400	108.1756	328	NDL / LE-3Dlite
5	Phuoc Hiep	PHI	15.4954	107.9776	56	NDL / LE-3Dlite
6	Tra Doc	TDVB	15.3342	108.1634	113	Guralp CMG-6TD (Q330 / Trillium-40)

The common seismic network was called VERIS (Vietnam Reservoir Induced Seismicity). Stations provided by IGFPAS were equipped with short-period seismometers Lennartz LE-3Dlite (1s), whereas stations provided by IGPVAST were equipped with long-period seismometers Guralp CMG-6TD (30s). Signals from seismometers Lennartz were recorded by Net Data Logger (NDL), which served in the project with a sampling rate of 100 samples per second (sps) and dynamics 132 dB. Seismometer Guralp has on-board digitizer with dynamic of 130 dB. Seismic signal is sampled with a frequency of 100 sps. Both systems are appropriate to measure local and regional seismicity, as the main content of seismic waves comes in the range of a few Hz, whereas stations Guralp were more suitable for the largest events in the ST2 region because of the lower frequency signal. The responses of both recording systems are shown in Fig. 2. Seismometers Lennartz 3Dlite were installed in new stations, whereas seismometers Guralp replaced the previously working recorders.

No.	Station name	Code	Co-ordination		Elevation [m]	Logger / Seismometer
			φ [°N]	λ [°E]		
7	Tra Bui	TBVB	15.3667	108.0503	224	Guralp CMG-6TD (SAMTAG / Guralp)
8	Tra Nu	TNVB	15.3564	108.3268	126	Guralp CMG-6TD (SAMTAG / Guralp)
9	Tra Mai	TMVB	15.1480	108.1202	202	Guralp CMG-6TD (SAMTAG / Guralp)
10	Tien Lanh	TLVB	15.4958	108.1200	325	Guralp CMG-6TD (SAMTAG / Guralp)

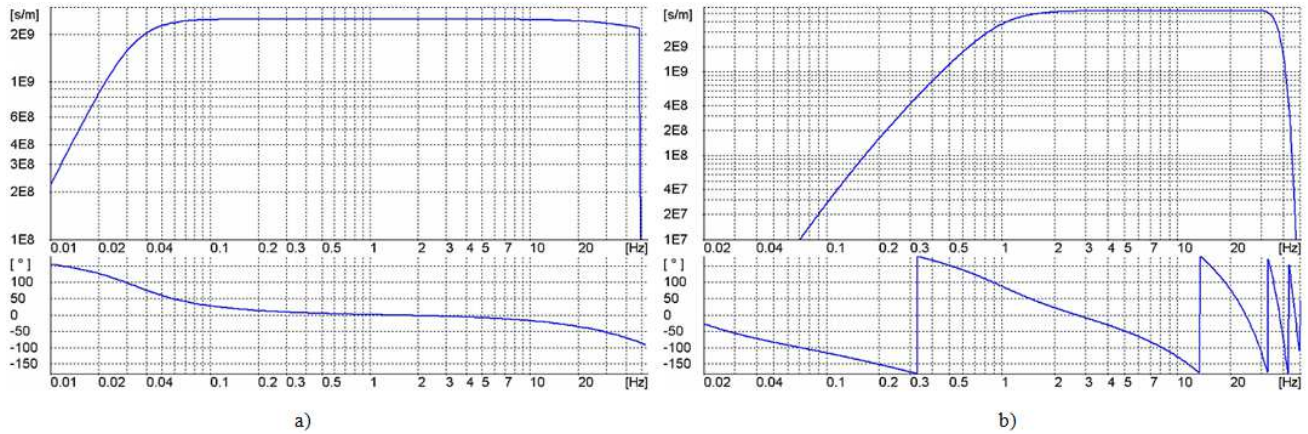
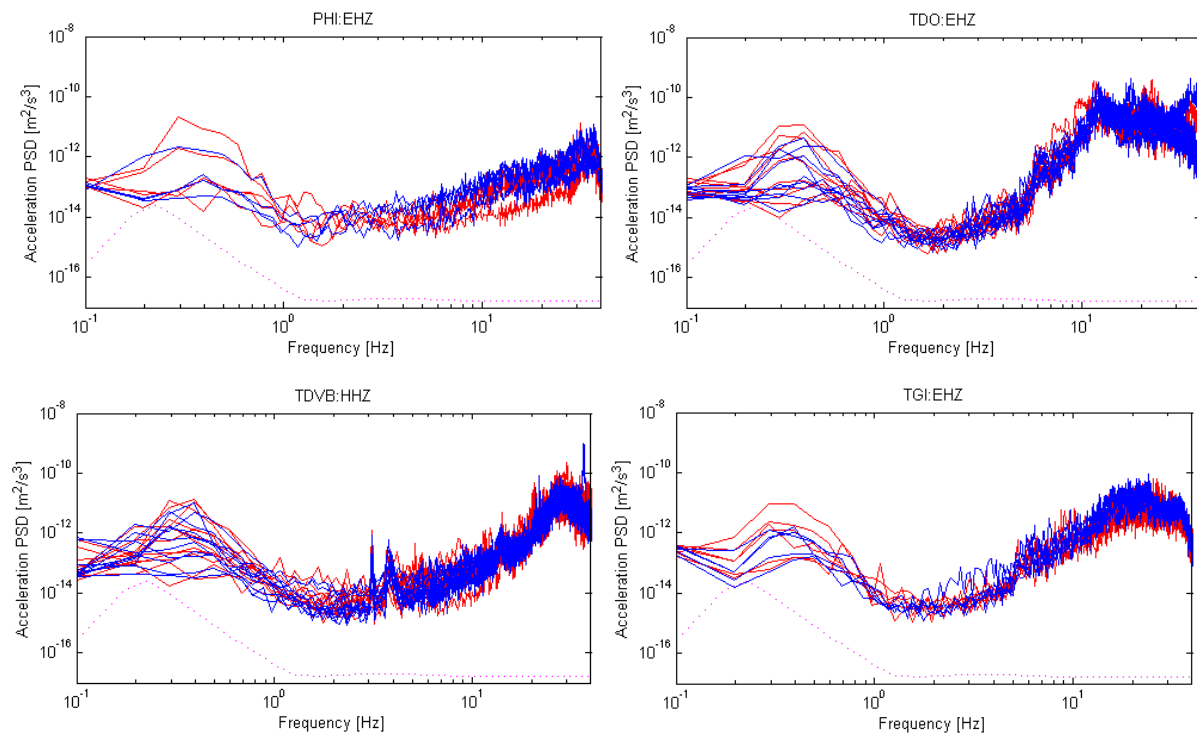


Fig. 2. The frequency- amplitude-phase response of: a) the Guralp CMG-6TD seismometer and b) the Lennartz LE-3Dlite seismometer with the NDL recorder.

The choice of location for installed local seismic stations in this region was preceded by Ground Penetrating Radar (GPR) measurements [10]. The goal of the investigation was to find the location and foundation for stations on the bedrock. The GPR measurement helped in assessing the cohesiveness of the sub-stratum, whether the rock is fissured or separated from the bedrock. Good foundation of a seismic stations allows noise reduction and enhances the early detection local seismic events. It is important, because there is the need for stations to be located in order to ensure security in neighborhoods. The basement investigation carried out by 100 MHz frequency antennas Pulse Ekko

100A Geological System for all 10 local seismic stations: Tra Doc, Tra Bui, Tra Nu, Tra Mai, Tra Don, Tra Leng, Tra Giac, Phuoc Hiep, Tien Lanh and Tien Ngoc. The interpretation of GPR radagrams showed good quality of data and near-surface geological structure consists of 4 layers where the three first layers are fractured from surface to depth of 3m and the last layer is slightly fractured and consolidated to 10m of depth. The Power Spectral Densities (PSD) of seismic noises are presented in Fig. 3. We didn't notice differences in noise between day and night and between dry and rainy seasons.



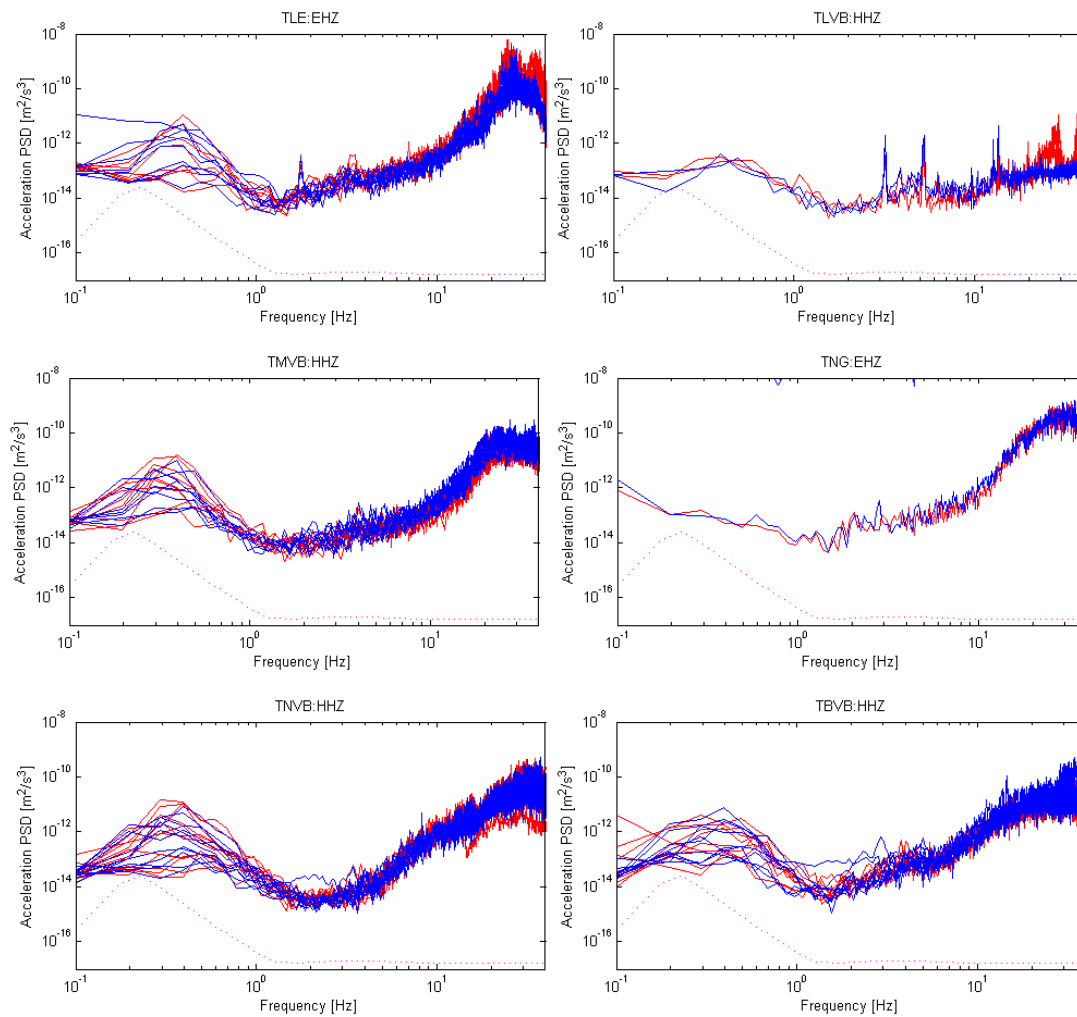


Fig. 3. The Power Spectral Densities (PSD) of seismic noises for local seismic stations of ST2 networks. In most stations the PSD is taken from both rainy and dry seasons. The electronic red lines describe PSD in day whereas blue ones show noise at night. The dotted line is the PSD of Peterson's (1993) Low Noise Model [18].

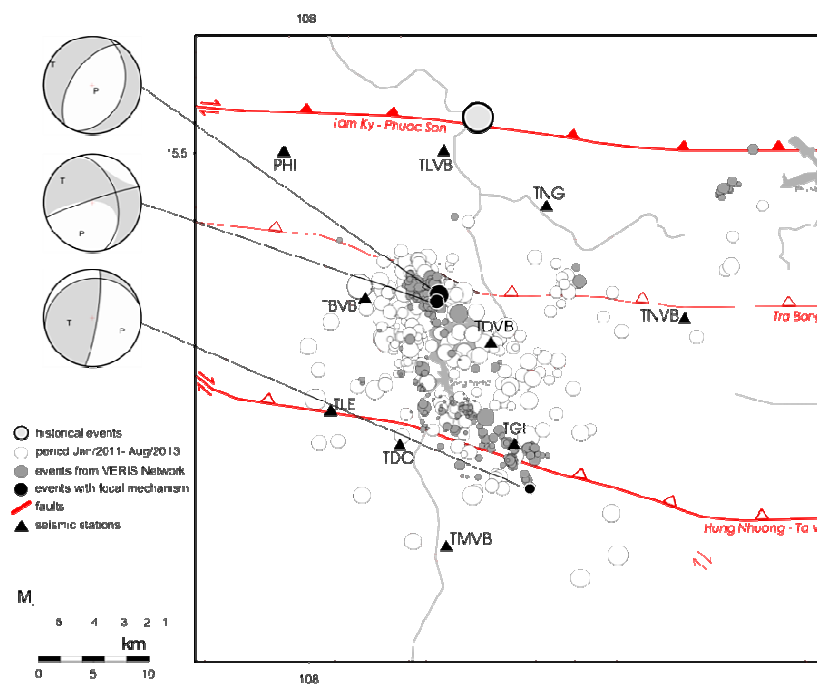


Fig. 4. Location of historical events and events from the period 2011-2014 grouped around the ST2 reservoir.

The final seismic network increased earlier possibilities of detection and location of weaker seismic earthquakes. It is also possible to calculate extended source parameters like focal mechanism, slip direction etc. Within the period from August 2013 to May 2014 about 2000 seismic events were detected, and 359 of them were localized and magnitudes were calculated (Fig. 4.). Seismic bulletins and catalogues of seismic monitoring in the ST2 region contain movements since August 2013. Seismic analysis and update catalogues are currently being conducted. The LocSAT application was used to locate events registered from the VERIS network. LocSAT with the IASP91 travel-time tables are used to perform hypocentral inversion of the phase arrival data. LocSAT estimates the origin time, epicentral location, and depth as an iterative least-squares inversion of travel time, slowness, and/or azimuth measurements [3,4].

Earthquakes registered in the period between January 2011 and July 2013 were of magnitude $1 \leq M_L \leq 4$. After the expansion of seismic networks in the ST2 region the threshold of completed recorded events were from $M_L = 0.3$. Earthquakes from August 2013 to May 2014 were recorded by the VERIS network in several places to the north-west and south-east from the ST2 reservoir. While these groupings were not visible from January 2011 to August 2013, because it was a larger location. We can specify two main groups of epicenters of earthquakes in the ST2 area, in the northern part of the reservoir near the Tra Bong thrust and the epicenters in the southern part arranged in a line sub-latitude to the likely Hung Nhuong – Ta Vi thrust.

4. Completed Catalogues for the Veris Network – Period III

Magnitude of completed (M_C) earthquake catalogues were the lowest magnitude at which all the earthquakes were detected in selected space and time volume. The completed catalogues were estimated on a frequency-magnitude distribution (FMD) of detected events in the selected region as well as by techniques [19]: Maximum Curvature (MAXC) [27], Entire Magnitude Range (EMR) [25,28] and study of the b-value stability of Gutenberg-Richter model [13] as a function of cut off magnitude (MBS) [5].

The Maximum Curvature method estimates M_C as the point of the maximum curvature by computing the maximum value of the first derivative of the frequency-magnitude curve (cumulated FMD). In practice, this matches the group of magnitudes with the highest frequency of events in the FMD. This technique requires fewer events than other techniques to reach a stable result, however it underestimates the M_C value sometimes [27].

The EMR technique is used to estimate the M_C the entire magnitude range with events below the M_C . Woessner and Wiemer (2005) [28] proposed a model consisting of two parts: the Gutenberg-Richter law for the complete part, and the cumulative normal distribution for the incomplete part of the

FMD.

Cao and Gao (2002) [5] estimated M_C using the stability of the b-value as a function of cut-off magnitude M_{C_0} , named as MBS by Woessner and Wiemer (2005) [28]. The M_C is defined as the magnitude for which the changes in b-value (Δb) is smaller than 0.03. It is based on an assumption that the estimation of b increases for $M_{C_0} < M_C$ and remains constant for $M_{C_0} > M_C$. This method doesn't produce good results in case of high variability of the FMD. Woessner and Wiemer (2005) [28] used the b-value uncertainty δb according to Shi and Bolt (1982) as criterion:

$$\delta b = 2.3b^2 \sqrt{\frac{\sum_{i=1}^N (M_i - \langle M \rangle)^2}{N(N-1)}} \quad (1)$$

where $\langle M \rangle$ is the mean magnitude and N is the number of events. In case of events in the Podhale region, the variability of the FMD is rather high, because of a small number of events.

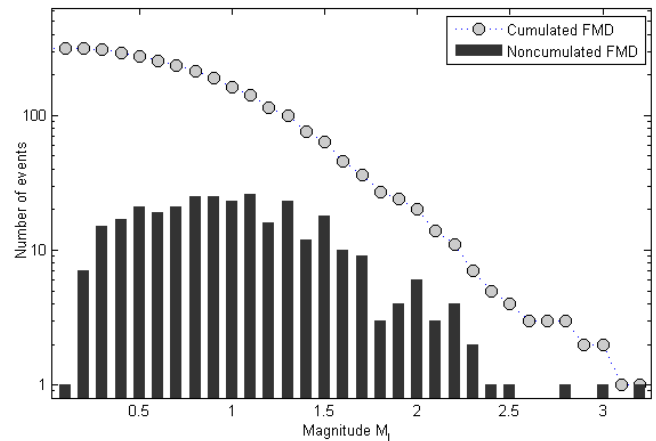


Fig. 5. The FMD of events recorded in ST2 region in period 2013 - 2014.

In the period from August 2013 to May 2014 VERIS seismic network in central Vietnam recorded about 2000 earthquakes where about 350 were located. The values of magnitudes M_C of completed catalogues were calculated by different methods and minimum magnitude (M_{min}) are presented in Tab. 3. Based on these values and based on the shape of the FMD of events recorded in years 2013-2014 the M_C was assumed as a value 1.1 (Fig. 5.) not dear more recorded events are less than M_C .

Tab. 3. Magnitudes of completed catalogues (M_C) of events in the period 2013-2014 estimated by various methods.

Technique	M_C
MAXC	0.8
EMR	1.1
MBS	0.22
M_{min}	0.12

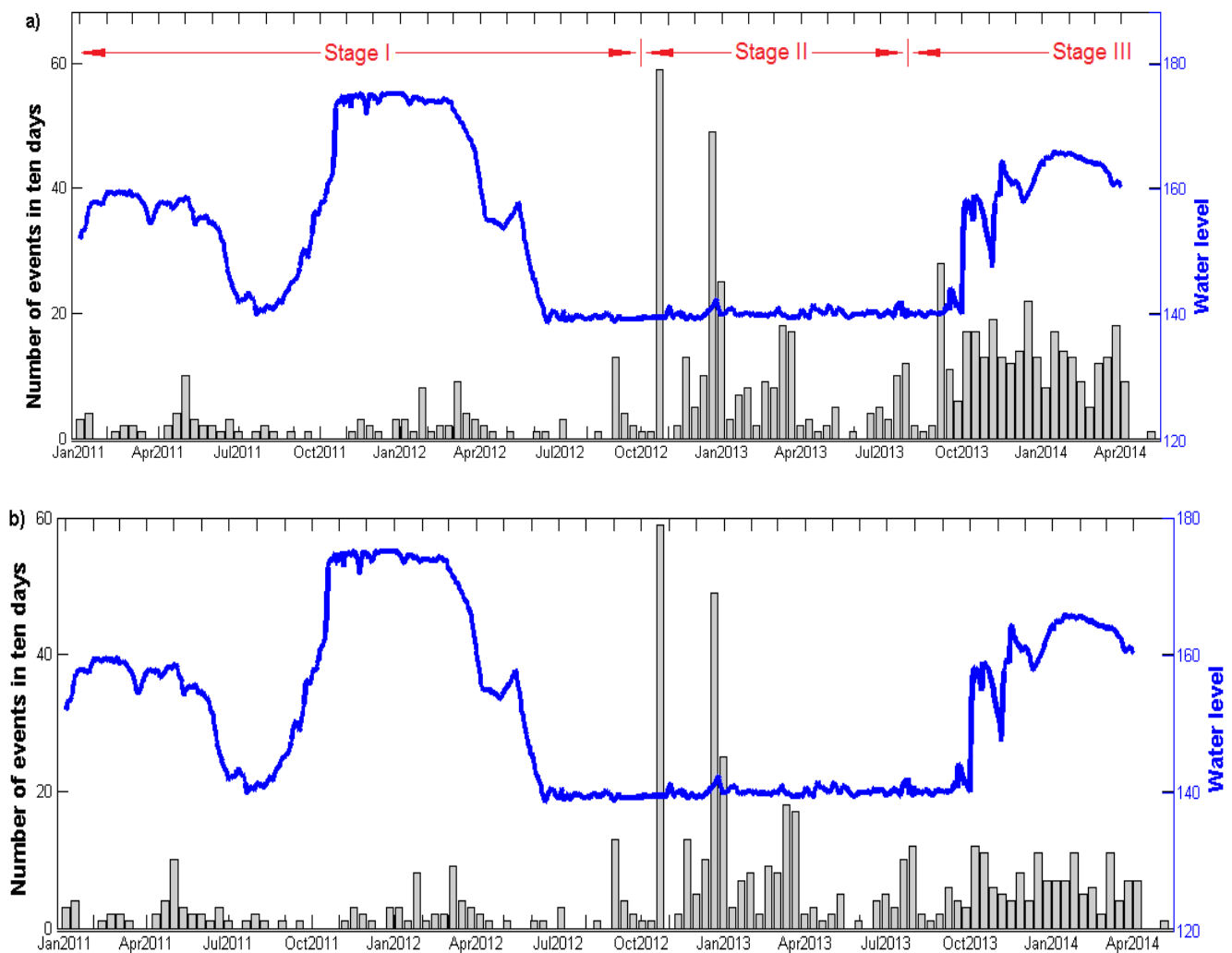
5. Seismic Activity Caused by the Filling of the ST2 Reservoir

The filling of the ST2 reservoir started in November 2010. The water level began to increase slowly from about 153 m in January 2011 to about 159.4 m in February 2011 and then it dropped to around 140 m in July 2011. The filling of the reservoir started again in August, 2011. In October, 2011, the maximum water level of 175 m was reached. It is also the highest level of reservoir activity till date. The water level of the reservoir was kept around 175 m until February 2012, when it was quickly dropped to the level of 157.64 m. In May, 2012, it was further decreased to 139.3 m. After which it was kept about 140 m to the end of September 2013. The water level began to increase with one sharp decline in November to the level of 149 m. The water level reached about 165.6 m in the beginning of 2014. Fig. 6 shows the relationship between the reservoir water level variation and earthquakes.

The natural seismic activity of the ST2 and its surroundings was very low. From the earthquake sources such as operational history, international data and data

recorded by seismic stations in Central Vietnam, it was found that between 1775 to 1992 only 13 eruptions were occurred in this area. However, there was an earthquake with magnitude of 4.7 in 1715 which was located near the ST2 reservoir (Fig.4) [24].

From the end of 2010, through analyses of earthquakes recorded by 2 Vietnam national seismic stations located in Binh Dinh and Hue (about 160-120 km far from ST2), it was found that several earthquakes appeared, whose epicenters were located in the province of Quang Nam. Until early 2011, the frequency of earthquakes in this area increased significantly. The occurrence of noticeable earthquakes in the ST2 area and its vicinity started in March, 2011 (Fig.6d). Especially, two quite strong eruptions occurred on October 22nd and November 15th, 2012, with magnitude $M_L=4.6$ and 4.7 respectively, which was felt by the inhabitants and damaged buildings in Bac Tra My area. On November 2012 the number of events in the 10-day period was high with a decreasing trend. Seismic activity is still in progress with an increase in the water level to 165 m.



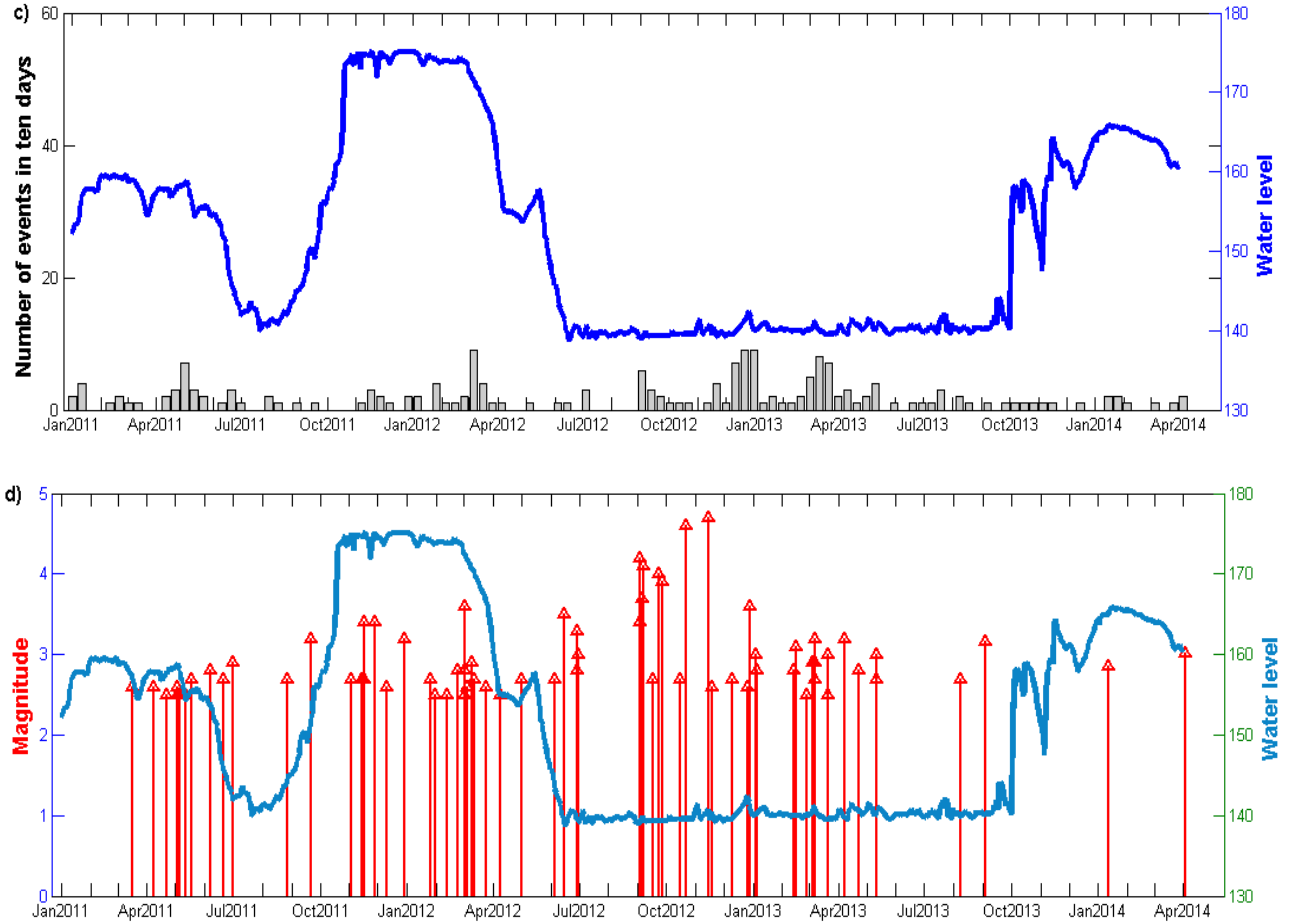


Fig. 6. The level of the water in the ST2 reservoir and seismic activity around the region in the period of January 2011 to April 2014 with: (a) number of all localized events in 10 days period, (b) number of events in 10 days period and magnitude higher than 1, (c) number of events in 10 days period and magnitude higher than 2, (d) events with magnitude higher than 2.5.

Fig. 6 a, b and c, shows that each stage increased earlier possibilities of detecting and locating smaller seismic earthquakes. These figures show improvement of seismic monitoring of the ST2 whereas Fig. 6 c and d shows the relationship between water level and seismicity for all the periods. The water level of reservoirs was tentatively divided into three cycles of high levels. During the first cycle when the water level reached the high of 160 m in February 2011, the frequency of earthquakes increased especially in times when the water level started falling in May 2011. This phenomenon also occurred in second cycles of high water from January 2011 to March 2012, when earthquakes began to appear the water level reached a high of 175 m with dense occurrence of events when the water level dropped from 173 m to 170 m in March 2012. During the third cycle from October 2013 seismic activity was relatively low, although the research continued. It suggests that seismicity is related to drawdown levels of the reservoir rather than to higher water levels like the lake in the case of Tehri Dam in India [8]. However, the highest density of earthquakes and the strongest eruptions appeared within the period from August 2012 to February 2013 when the water level was low. Additionally, the decision to reduce the level of water is often caused by increasing the seismicity.

6. The Focal Mechanism of Events: September 3rd, 2013 with $M_1 = 3.2$, October 24th, 2013 with $M_1 = 2.3$, October 31st, 2013 with $M_1 = 1.8$

Calculations of the moment tensor from the VERIS network records were performed by FOCI software 3.0 [17]. The inversion of the P wave amplitude was performed in time domain approach [2, 26]. The registered first onsets were: Direct P wave according to the velocity model of the ST2 area. The velocities determined for the waves were: Direct wave 5 km/s. The Direct waves were recorded at 8 stations on the VERIS network. The input parameters are the amplitude and polarity information on the first P-wave displacement pulses. According to Fitch et al. (1980) [9] the recorded displacement for the vertical component of the P-wave phase is:

$$U_z^P(x, t) = \frac{1}{4\pi\rho\alpha^3 r} \left[\bar{y} M \dot{s} \left(t - \frac{r}{\alpha} \right) \bar{y} \right] l_z, \quad (2)$$

Where ρ is the average density, r is the source-receiver distance, α is the average velocity of P wave, M is the seismic moment, l_z is the cosine of the angle of the incidence and is the take-off angle. The Source Time Function (STF) was based on the Haskell's source model [15]:

$$\dot{s} = \begin{cases} 1/T, & 0 < t < T \\ 0, & \text{elsewhere} \end{cases} \quad (3)$$

where T is rupture time.

The moment tensor is obtained by solution of a set of N equations of type (1). From Awad and Kwiitek, 2005 [2], Cesca et al., 2010 [7], we have $U=GM$, where M is moment tensor and G is Green function, U is the displacement matrix. Inversion of this formula by solving it with use of the (1) allows to find moment tensor. Determination of the Moment tensor is realized by finding the best solution of $(GM-U)^2=\min.$, and U is the displacement amplitude calculated with eq. (1).

Due to the six independent elements of the moment tensor there must be at least six such equations, but the more, the better. The deviatoric, pure shear and full moment tensor was calculated using the L_2 norm as a measure of the misfit [2, 26]. The results of inversion are presented in Tab. 4 and Fig. 7. The focal sphere azimuthal coverage was sparse and number of available recordings was at the limit of the methodological requirements, which obviously influenced the inversion results. Nevertheless the first three focal mechanism solutions are a good first insight into the focal and tectonic properties of the process inducing seismic events in the vicinity of the ST2.

Source mechanism of September 3rd, 2013 event is a normal fault with strike azimuth NNE-SSE and moderate dip angle about 40-50 degrees (Fig.7). Nodal plane orientations are sub-parallel to the Tra Bong thrusting movement direction, but perpendicular to the strike of this main

discontinuity. In the area of the event location the main discontinuities are not in agreement with the nodal planes' orientation. Moreover, the main stress regime indicates strike-slip faulting, while the inversion results indicate the normal faulting. The tension axis orientation are similar to the main strike slip stress pattern in this area, but the polarities and amplitudes of the first P-wave pulses determine the focal mechanism results as a normal fault. The quality of the solution is limited by the number of stations available and their azimuthal coverage. There are only 8 stations with good signal to noise ratio, which were used in inversion. The azimuthal coverage was good in the southern half of the focal sphere, but only one station was available in the northern half of the focal sphere.

Source mechanism of October 24th, 2013 event is a strike-slip fault with a NEE-SWW or NNW-SSE strike. The latter strike orientation is parallel to the local fault connecting the two main discontinuities Tra Bong and Hung Nhuong- Ta Vi. The azimuthal station coverage of the focal sphere was poor in the NW quadrant and reasonably well in the southern half. The event was located near the location of the September 3rd, 2013 eruption, but their focal mechanism and fault plane solutions differ.

Source mechanism obtained for the October 31st, 2013 event was a reverse fault with strike almost N-S or W-E. The second nodal plane orientation is almost similar to the Hung Nhuong -Ta Vi thrust, but the main regime of this area is strike-slip, which is not in agreement with the obtained focal mechanism. Again the azimuthal station coverage was not good, especially in the SW quadrant of the focal sphere, where even a single station was not available, while 6 out of 8 available stations covered the NW quadrant.

Obtained results are not sufficient for any general conclusions about the stress or tectonic regime of the process leading to the seismicity of the ST2 reservoir vicinity.

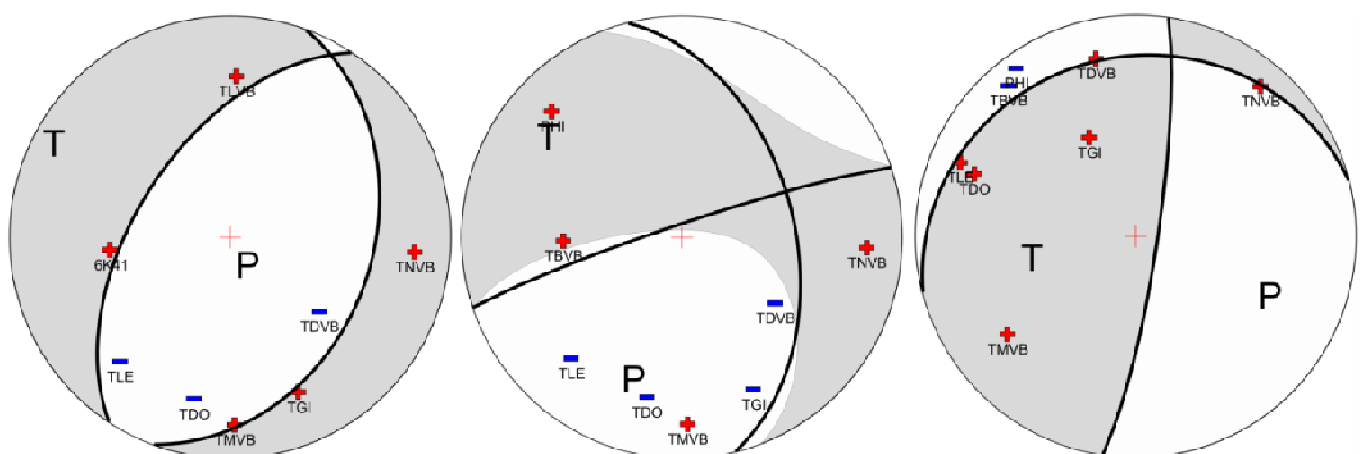


Fig. 7. Mechanism solutions with P-wave amplitude inversion of three events from the left: September 3rd, 2013 $M_L = 3.2$, October 24th, 2013 $M_L = 2.3$, October 31st, 2013 $M_L = 1.8$. The codes and signs denote the station location and polarization on focal sphere. All solutions are presented as a Lambert azimuthal equal-area projection on upper hemisphere Schmidt net.

Tab. 4. Seismic Moment, Moment Magnitude and nodal planes of the three seismic events calculated from VERIS network.

Event	Seismic Moment [Nm]	M _w	Nodal plane A	Nodal plane B
September 3 rd , 2013	$5.01 \cdot 10^{13}$	3.1	33°/55°/-82°	200°/36°/-101°
October 24 th , 2013	$1.44 \cdot 10^{12}$	2.1	76°/89°/-51°	167°/39°/-179°
October 31 st , 2013	$2.85 \cdot 10^{12}$	2.2	190°/78°/72°	69°/22°/147°

7. Conclusions

Before 2010, when ST2 reservoir was not yet in operation, earthquake activity in the area was low, but from November 2010, when the reservoir was full of water, the earthquake occurrence increased, most notable were the large two quakes 4.6 and 4.7 Richter on November 2012. This event is the basis to link past earthquake-causing sources related to reservoir stimulation earthquake here. The seismicity is clearly triggered by the ST2 reservoir as the region had been practically aseismic before the dam was built.

The characteristics of earthquake activity stimulated reservoir area are the initial results based on earthquake records data string is not complete by the early use of the reservoir. Since the local earthquake recording stations network was put into operation, the data source record was able to really handle, analyze which results show seriousness of earthquakes in this area concentrated in the protection system of the fault Tra Bong and Hung Nhuong-Ta Vi along in the direction of NNW-EES and the cause of the increased activities of the earthquakes are related with respect to the operation of the ST2 reservoir. It also suggests that reservoir construction is a major factor in the seismogenic process.

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